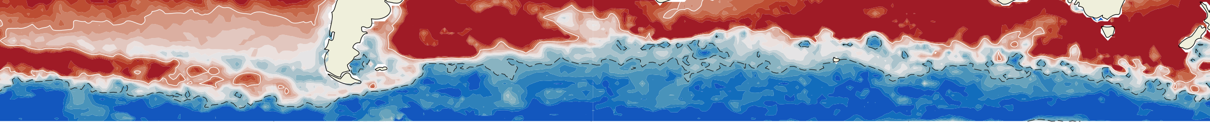


From alpha to beta ocean

Exploring the role of surface buoyancy fluxes and seawater thermal expansion in setting the upper ocean stratification

Romain Caneill

Seminar ENSL, March 19, 2024

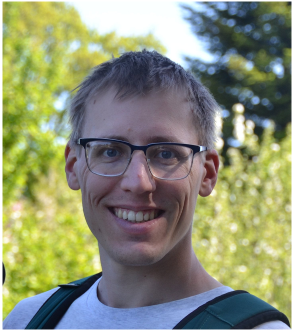
PhD Supervisor: Fabien Roquet
Collaborators: Gurvan Madec and Jonas Nycander

2014 – 2017 ENSL, physics and geophysics

2017 – 2018 CAP de menuiserie

2018 – 2024 PhD, Göteborg with Fabien Roquet

2024 – 2027 Postdoc, Grenoble, SASIP project



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PGP fingerprint:

70D5 7116 37B2 9335 9088

F124 D0FE 114E BFFD ED7F

Some of my interests:

- Python
- NEMO
- xnemogcm, xgcm
- Reproducible science
- Snakemake
- Free Software

For the science, follow this presentation :)

The oceans store carbon and heat

The oceans have taken up about:

- 25 % of CO_2 produced by human activities;

The oceans store carbon and heat

The oceans have taken up about:

- 25 % of CO_2 produced by human activities;
- 90 % of excess heat.

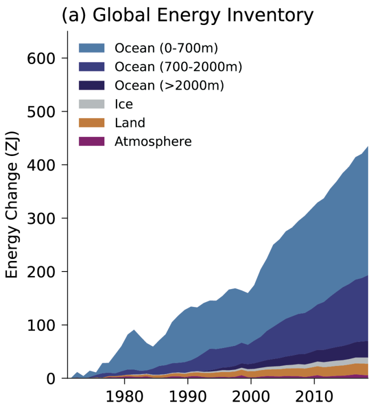


Figure adapted from the IPCC Sixth Report (Fox-Kemper et al., 2021)

Ocean and atmosphere exchanges properties through the mixed layer

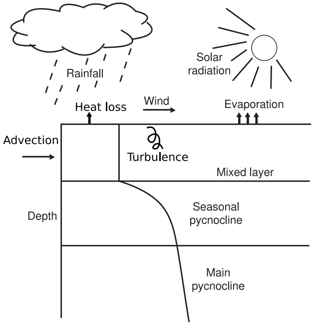


Figure adapted from Sprintall and Cronin (2009)

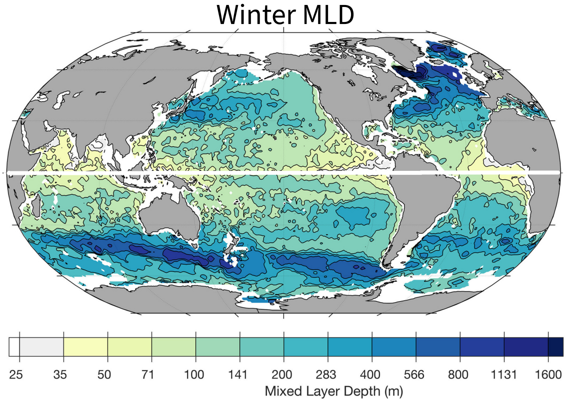


Figure adapted from Johnson and Lyman (2022)

The ocean absorbs anthropogenic CO_2

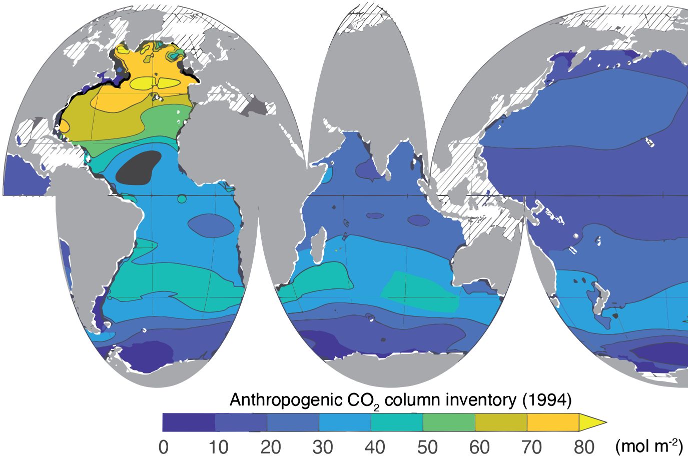


Figure adapted from Gruber et al. (2023)

The global circulation brings the water properties at depth

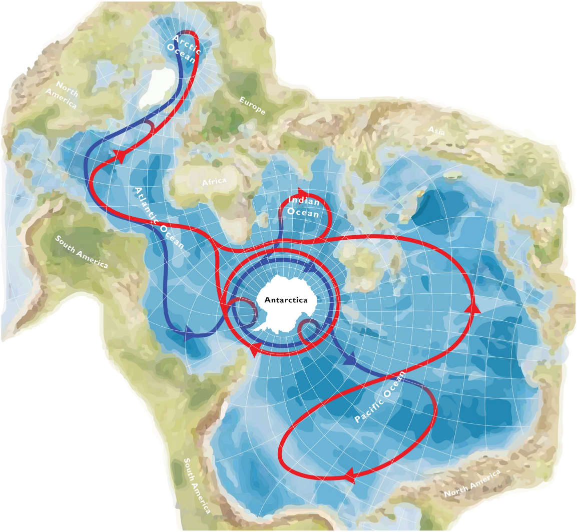
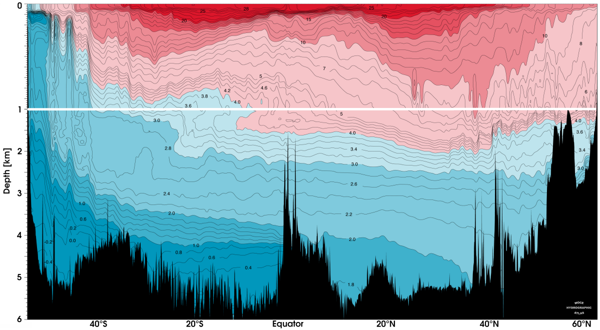


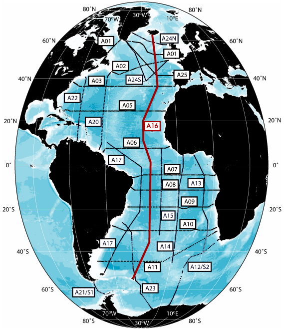
Figure adapted from Meredith (2019)

Ocean stratification

WOCE A16 section of potential temperature

The large stratification inhibits vertical exchanges.



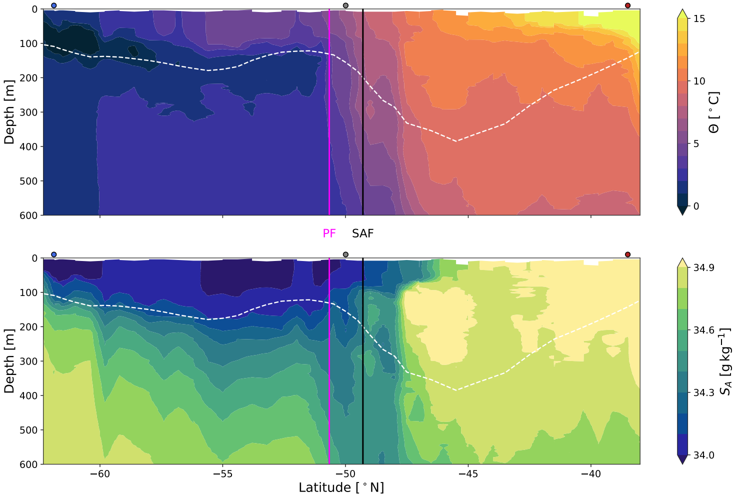


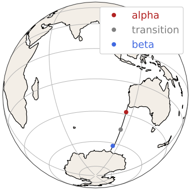
The ocean is mainly stratified because it is heated up at the surface.

Regimes of stratification

T-S section IO9S

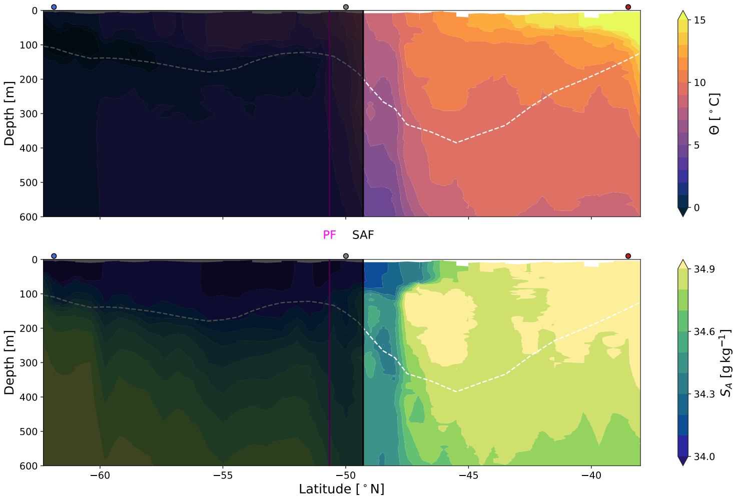
https://cchdo.ucsd.edu/cruise/09AR20120105

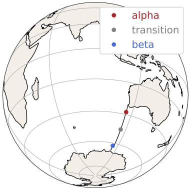


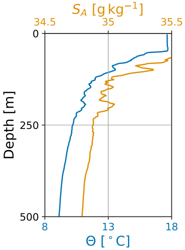


Regimes of stratification

Alpha ocean

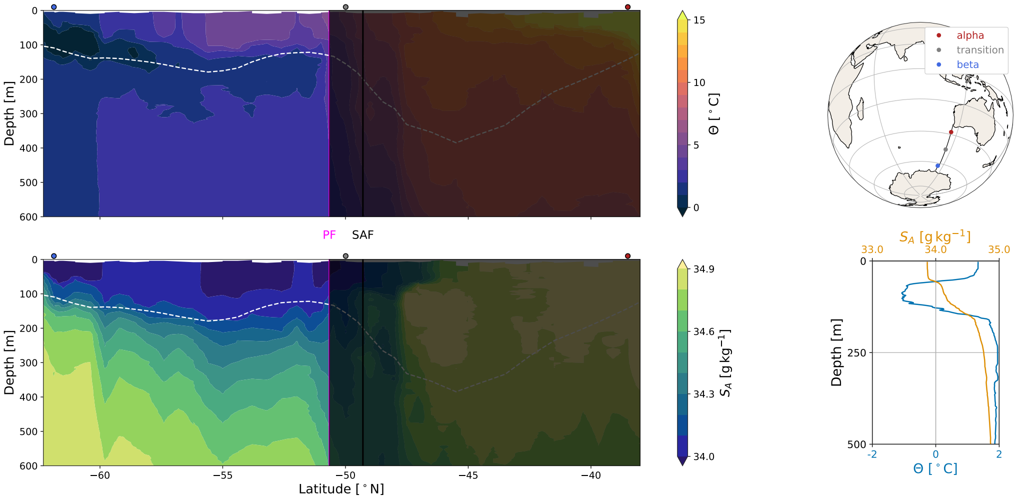






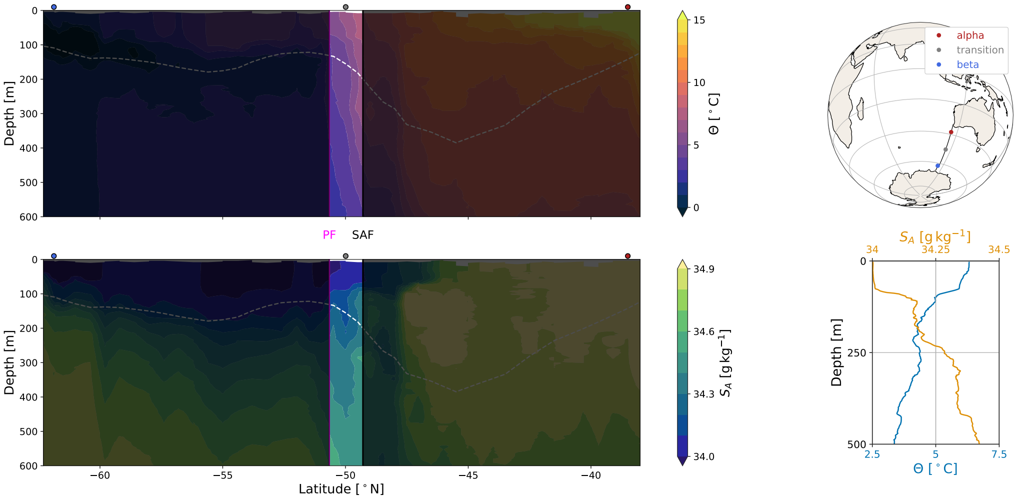
Regimes of stratification

Beta ocean



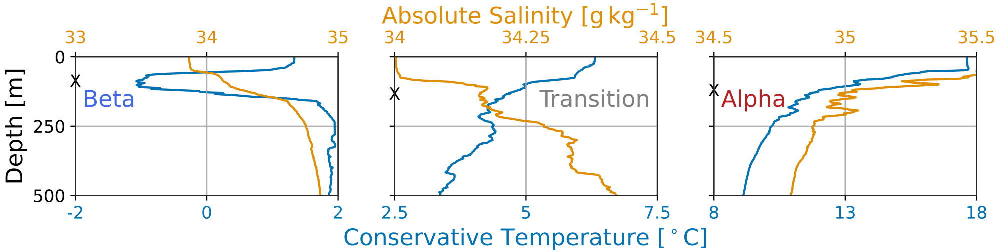
Regimes of stratification

Transition zone



Beta, transition, and alpha

T-S section IO9S, selected profiles

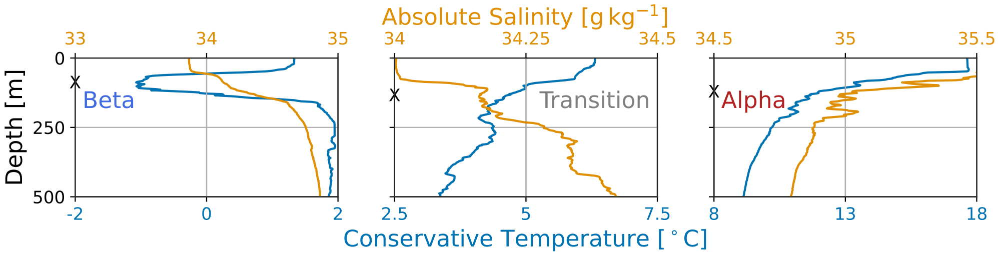


Temperature stratifies:
alpha ocean

(Carmack, 2007)

Beta, transition, and alpha

T-S section IO9S, selected profiles



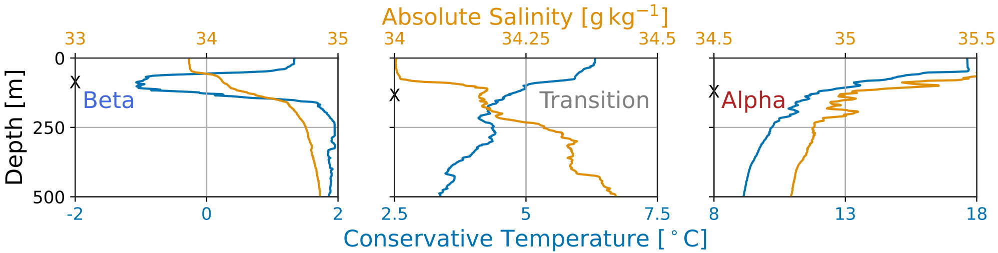
Salinity stratifies:
beta ocean

Temperature stratifies:
alpha ocean

(Carmack, 2007)

Beta, transition, and alpha

T-S section IO9S, selected profiles



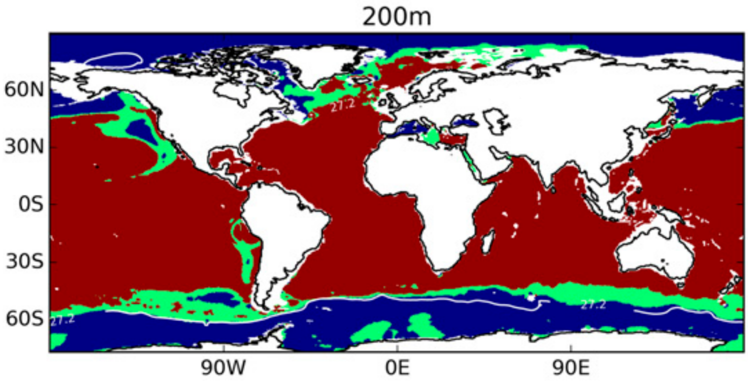
Salinity stratifies:
beta ocean

Both stratify:
transition zone

Temperature stratifies:
alpha ocean

(Carmack, 2007)

Alpha and beta oceans

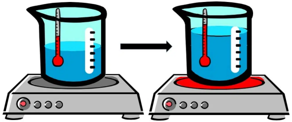


Called alpha – beta oceans in reference to α and β, thermodynamic properties of seawater.

Figure adapted from Stewart and Haine (2016)

The thermal expansion coefficient (TEC, α)

- Cold water is usually denser than warm water.



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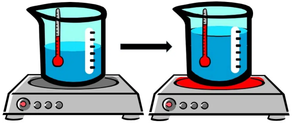
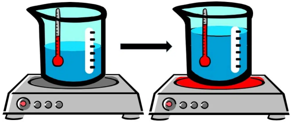
- Cold water is usually denser than warm water.
- Ocean warms $\Longrightarrow$ volume increases (1/2 of observed sea-level rise)



The thermal expansion coefficient (TEC, α)

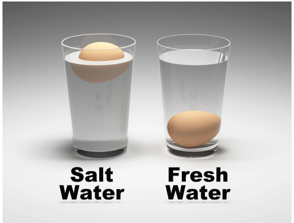
- Cold water is usually denser than warm water.
- Ocean warms $\implies$ volume increases (1/2 of observed sea-level rise)
- The TEC quantifies the relative change of density with temperature:

$$\alpha = -\frac{1}{\rho}\frac{\partial \rho}{\partial \Theta}$$



The haline contraction coefficient (HCC, β)

- Salty water is denser than freshwater



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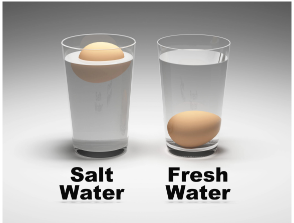
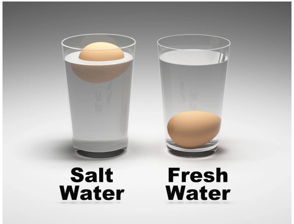


The haline contraction coefficient (HCC, β)

- Salty water is denser than freshwater
- The HCC quantifies the relative change of density with salinity:

$$\beta = \frac{1}{\rho}\frac{\partial \rho}{\partial S}$$



Properties of the TEC and HCC

- The TEC follows a (quasi) linear relation with temperature

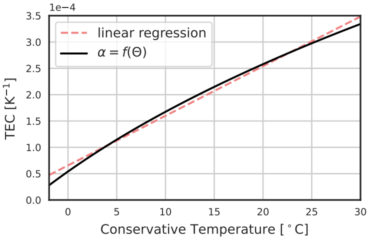


Figure adapted from Caneill et al. (2023)

Properties of the TEC and HCC

- The TEC follows a (quasi) linear relation with temperature
- The HCC variations in the ocean are negligible
  $\beta \simeq 7.5 \times 10^{-4}\,\mathrm{kg\,g^{-1}}$

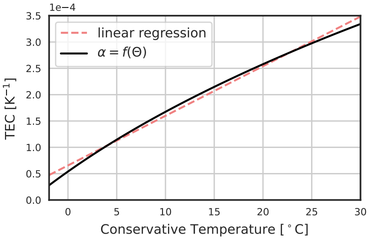


Figure adapted from Caneill et al. (2023)

Properties of the TEC and HCC

- The TEC follows a (quasi) linear relation with temperature
- The HCC variations in the ocean are negligible $\beta \simeq 7.5 \times 10^{-4}$ kg g^{-1}
- It was assumed that the role of salinity is enhanced in polar regions due to low values of the TEC

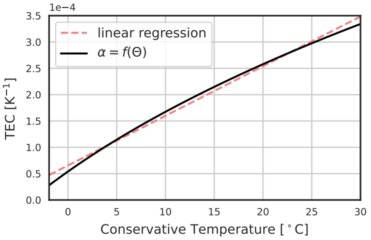


Figure adapted from Caneill et al. (2023)

What is the origin of alpha and beta oceans?

Objectives

From alpha to beta ocean: Exploring the role of surface buoyancy fluxes and seawater thermal expansion in setting the upper ocean stratification

Objective A	Objective B	Objective C
Describe alpha – beta oceans using observations	How do buoyancy fluxes shape the upper stratification?	Assess the role of the local value of the TEC.

TEC = Thermal expansion coefficient

Objectives

From alpha to beta ocean: Exploring the role of surface buoyancy fluxes and seawater thermal expansion in setting the upper ocean stratification

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How do buoyancy fluxes shape the upper stratification?

Objective C

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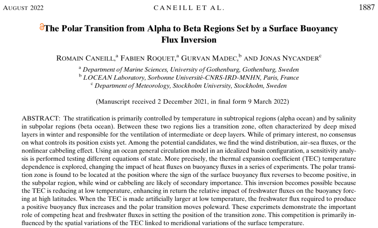
This presentation

Paper I

Caneill, R., Roquet, F., Madec, G., & Nycander, J. (2022). The Polar Transition from Alpha to Beta Regions Set by a Surface Buoyancy Flux Inversion. *Journal of Physical Oceanography*



AUGUST 2022 CANEILL ET AL. 1887

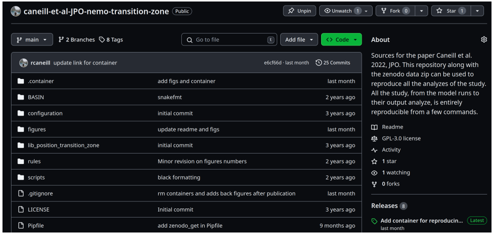
The Polar Transition from Alpha to Beta Regions Set by a Surface Buoyancy Flux Inversion

ROMAIN CANEILL,[a] FABIEN ROQUET,[a] GURVAN MADEC,[b] AND JONAS NYCANDER[c]

[a] *Department of Marine Sciences, University of Gothenburg, Gothenburg, Sweden*
[b] *LOCEAN Laboratory, Sorbonne Université-CNRS-IRD-MNHN, Paris, France*
[c] *Department of Meteorology, Stockholm University, Stockholm, Sweden*

(Manuscript received 2 December 2021, in final form 9 March 2022)

ABSTRACT: The stratification is primarily controlled by temperature in subtropical regions (alpha ocean) and by salinity in subpolar regions (beta ocean). Between these two regions lies a transition zone, often characterized by deep mixed layers in winter and responsible for the ventilation of intermediate or deep layers. While of primary interest, no consensus on what controls its position exists yet. Among the potential candidates, we find the wind distribution, air–sea fluxes, or the nonlinear cabbeling effect. Using an ocean general circulation model in an idealized basin configuration, a sensitivity analysis is performed testing different equations of state. More precisely, the thermal expansion coefficient (TEC) temperature dependence is explored, changing the impact of heat fluxes on buoyancy fluxes in a series of experiments. The polar transition zone is found to be located at the position where the sign of the surface buoyancy flux reverses to become positive, in the subpolar region, while wind or cabbeling are likely of secondary importance. This inversion becomes possible because the TEC is reducing at low temperature, enhancing in return the relative impact of freshwater fluxes on the buoyancy forcing at high latitudes. When the TEC is made artificially larger at low temperature, the freshwater flux required to produce a positive buoyancy flux increases and the polar transition moves poleward. These experimets demonstrate the important role of competing heat and freshwater fluxes in setting the position of the transition zone. This competition is primarily influenced by the spatial variations of the TEC linked to meridional variations of the surface temperature.



100 % reproducible with few commands

https://doi.org/10.1175/JPO-D-21-0295.1

https://github.com/rcaneill/caneill-et-al-JPO-nemo-transition-zone

This presentation

Paper II

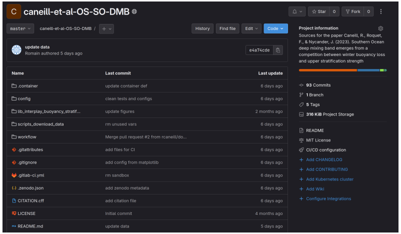
Caneill, R., Roquet, F., & Nycander, J. (2023). Southern Ocean deep mixing band emerges from competition between winter buoyancy loss and stratification. *accepted for publication in Ocean Science*

https://doi.org/10.5194/egusphere-2023-2404
Preprint. Discussion started: 19 October 2023

EGUsphere
Preprint repository

Southern Ocean deep mixing band emerges from a competition between winter buoyancy loss and upper stratification strength

Romain Caneill[1], Fabien Roquet[1], and Jonas Nycander[2]
[1]Department of Marine Sciences, University of Gothenburg, Göteborg, Sweden
[2]Department of Meteorology, Stockholm University, Stockholm, Sweden
Correspondence: Romain Caneill (romain.caneill@gu.se)



100 % reproducible with few commands

https://doi.org/10.5194/egusphere-2023-2404

https://gitlab.com/rcaneill/caneill-et-al-OS-SO-DMB

This presentation

Paper III

Caneill, R., & Roquet, F. (2023). Temperature versus salinity: Distribution of stratification control in the global ocean. *in preparation for Ocean Science*

This presentation

Paper IV

Roquet, F., Ferreira, D., Caneill, R., Schlesinger, D., & Madec, G. (2022). Unique thermal expansion properties of water key to the formation of sea ice on Earth. *Science Advances*

ScienceAdvances

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HOME > SCIENCE ADVANCES > VOL. 8, NO. 46 > UNIQUE THERMAL EXPANSION PROPERTIES OF WATER KEY TO THE FORMATION OF SEA ICE ON EARTH

RESEARCH ARTICLE | CLIMATOLOGY

Unique thermal expansion properties of water key to the formation of sea ice on Earth

FABIEN ROQUET, DAVID FERREIRA, ROMAIN CANEILL, DANIEL SCHLESINGER, AND GURVAN MADEC Authors Info & Affiliations

SCIENCE ADVANCES • 16 Nov 2022 • Vol 8, Issue 46 • DOI: 10.1126/sciadv.abq0793

Objective A	Objective B	Objective C
Describe alpha – beta oceans using observations	How do buoyancy fluxes shape the upper stratification?	Assess the role of the local value of the TEC.
Paper III	Papers I, II	Papers I, II, and IV

Paper III Caneill, R., & Roquet, F. (2023). Temperature versus salinity: Distribution of stratification control in the global ocean. *in preparation for Ocean Science*

Stratification Control Index (SCI)

$$SCI = \frac{\alpha \partial_z \Theta + \beta \partial_z S}{\alpha \partial_z \Theta - \beta \partial_z S} \tag{1}$$

The SCI quantifies the relative effect of temperature and salinity on stratification.

SCI > 1: alpha   -1 < SCI < 1: transition   SCI < -1: beta

Compute climatology of winter SCI

- Based on about 20 years of observation profiles (EN4 database, includes ARGO, ship-based CTD, MEOP, etc)
- Compute the SCI at the bottom of winter mixed layer
- Interpolation to produce climatology

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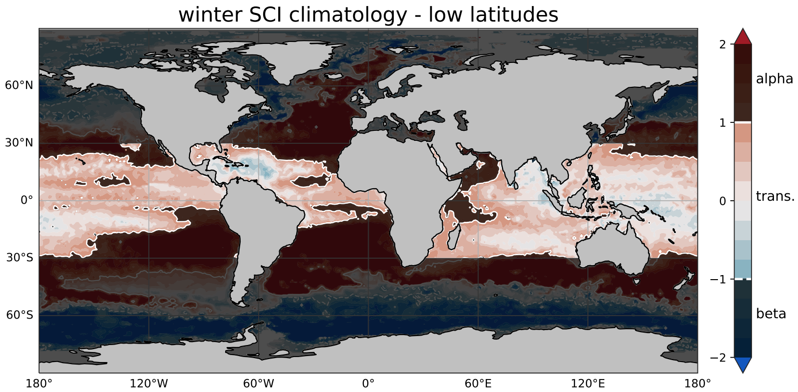
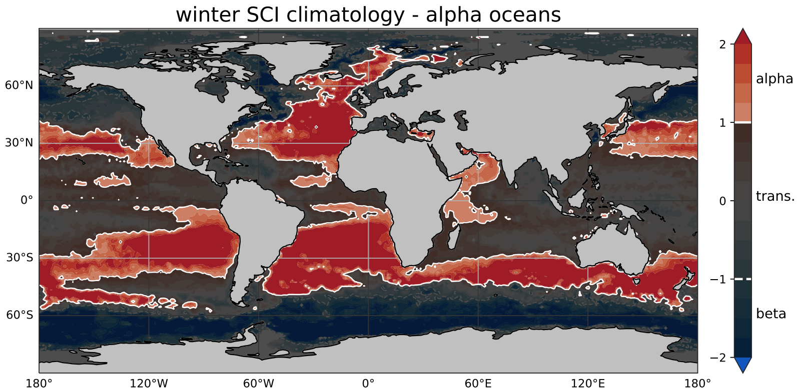


- Low-latitudes: transition zone
- Mid-latitudes: alpha ocean
- Between alpha and beta: PTZ
- High-latitudes: beta ocean

PTZ = polar transition zone

Global maps of the winter SCI (Paper III)



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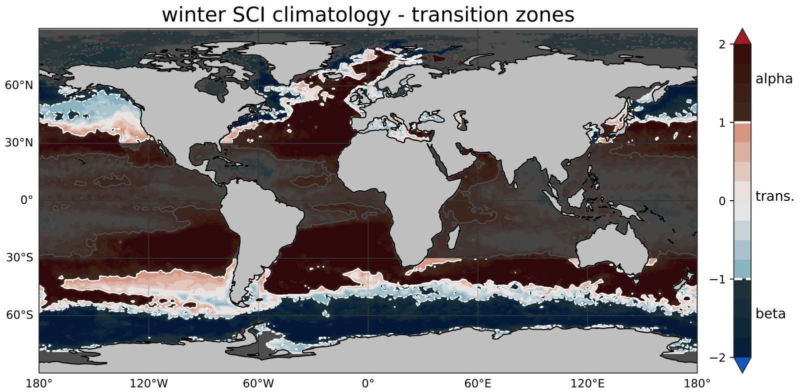

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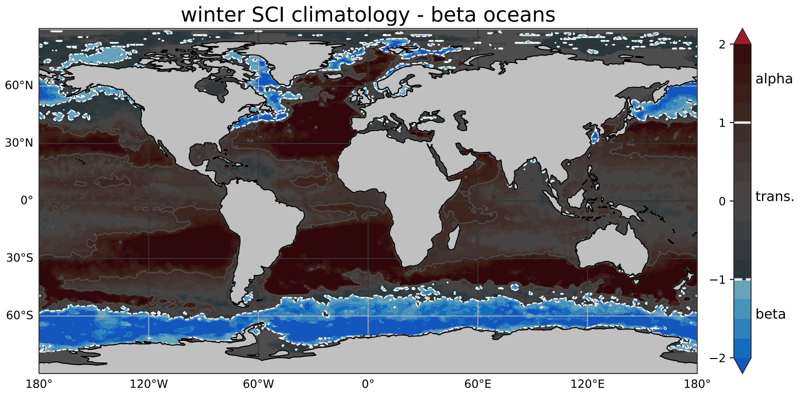

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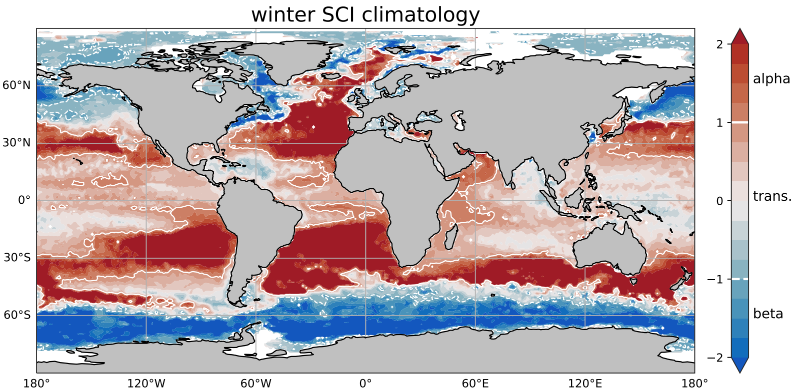
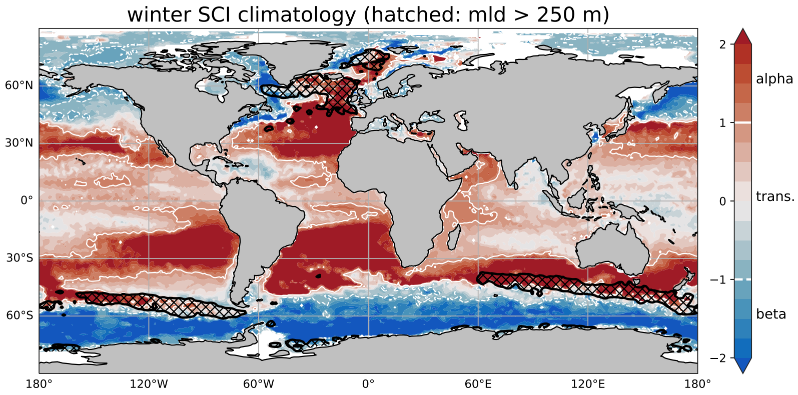


- Zonation with: transition zone → alpha → PTZ → beta
- Wide and zonal North Pacific PTZ
- Narrow and diagonal North Atlantic PTZ

Global maps of the winter SCI (Paper III)



Deep MLs located at the poleward flank of alpha oceans.

- Deep MLs mostly found in alpha oceans
- Bimodal distribution of the SCI, centred around ± 1.5

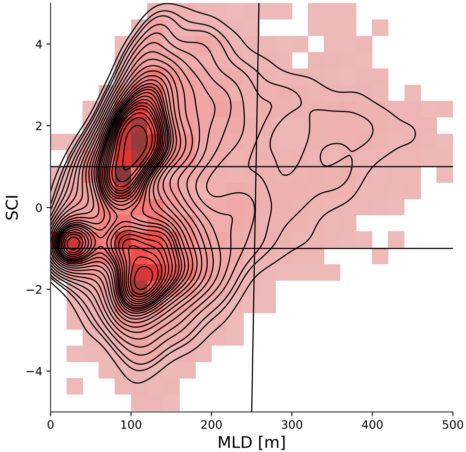


Figure for $|\varphi| \geq 30^{\circ}$

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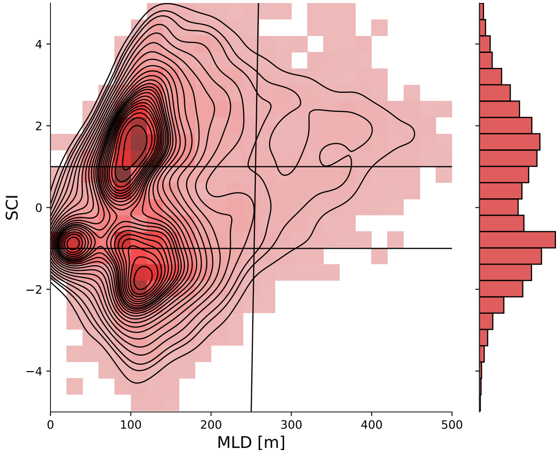

Figure for $|\varphi| \geq 30^{\circ}$

Objective B

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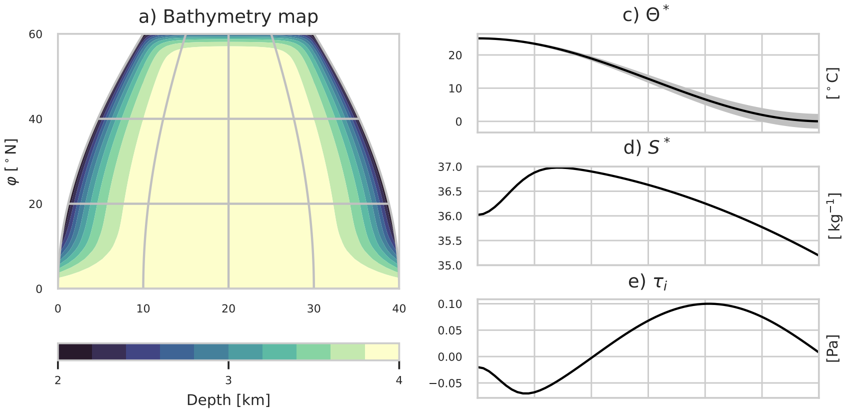
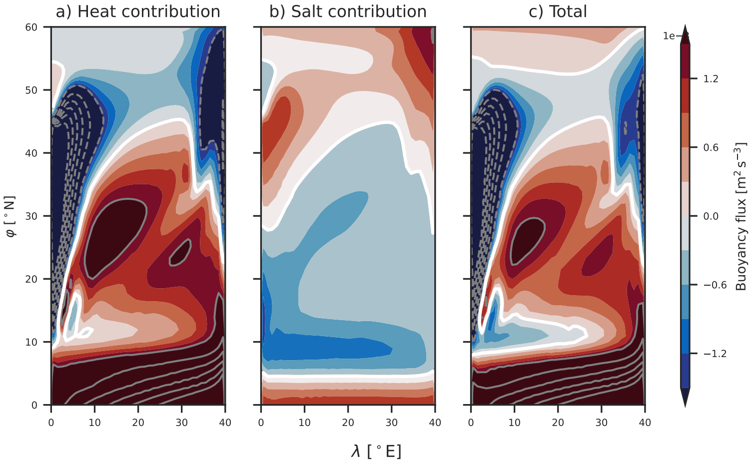
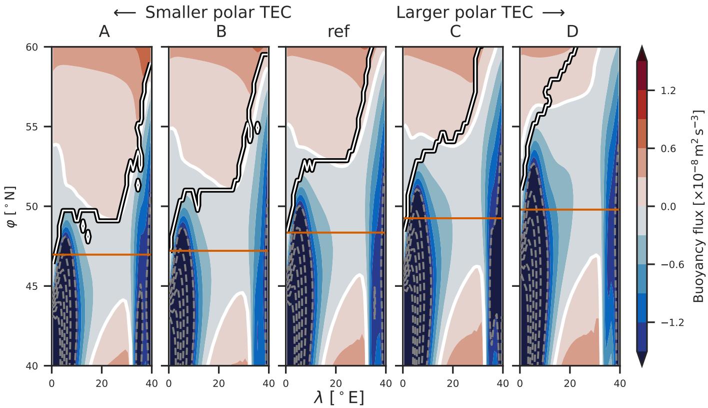


Idealised configuration that allows to study the role of annual buoyancy fluxes, by modification of the equation of state (thus changing the TEC).

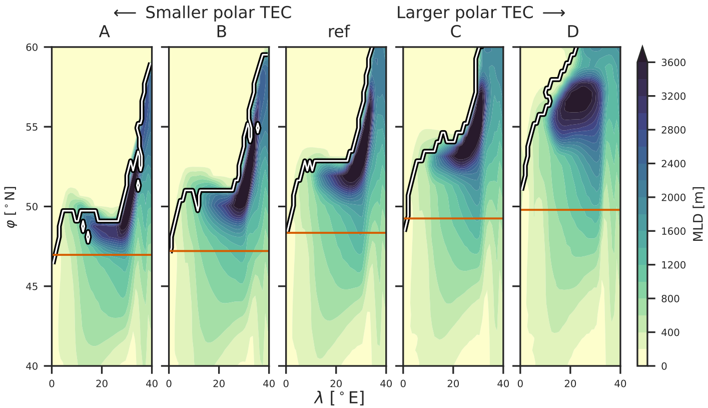


Reference run

Annual buoyancy fluxes set the transition (Paper I)



Wind kept unchanged!



Will fronts move poleward due to increased ocean temperature?

Poleward migration of transition zone due to global warming?

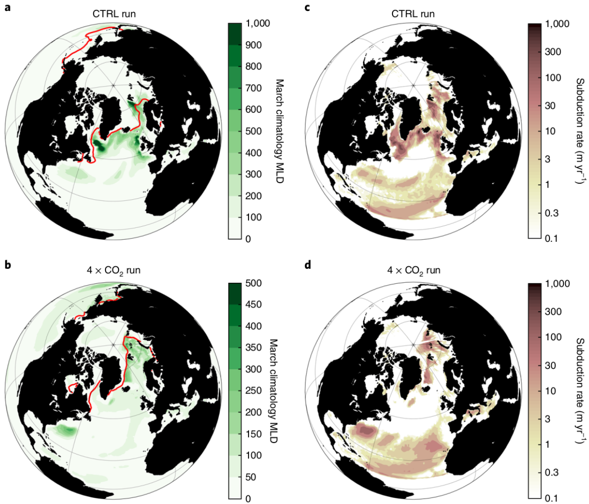
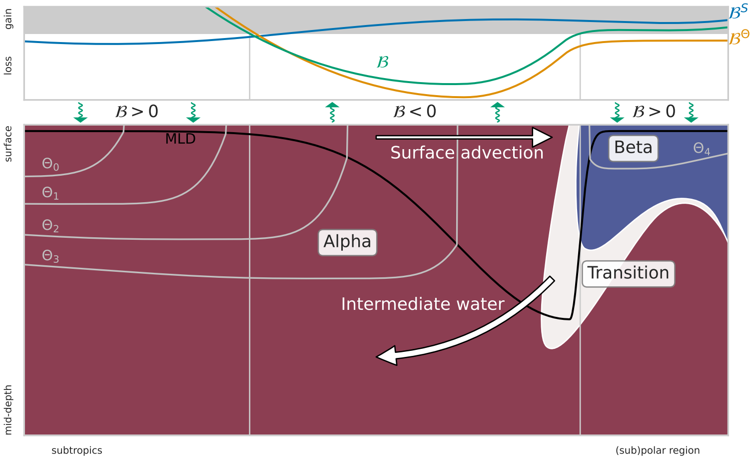
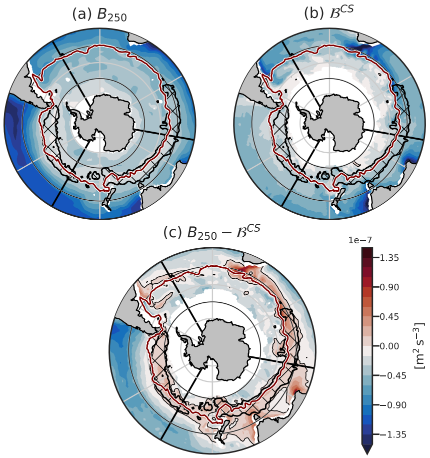


Figure from Lique and Thomas (2018)

Annual buoyancy fluxes set the transition (Paper I)



Winter buoyancy loss erodes stratification (Paper II)



- B_{250}: measure of stratification
- $\mathcal{B}^{CS}$: buoyancy loss
- Hatched region: the DMB

- The position of the deep MLs is set by the balance between buoyancy loss and stratification
- Buoyancy fluxes control the stratification regimes

DMB = deep mixing band

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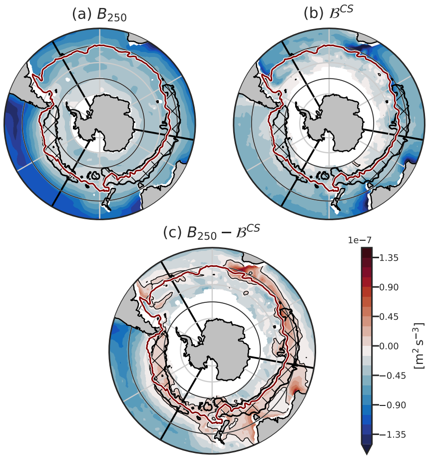

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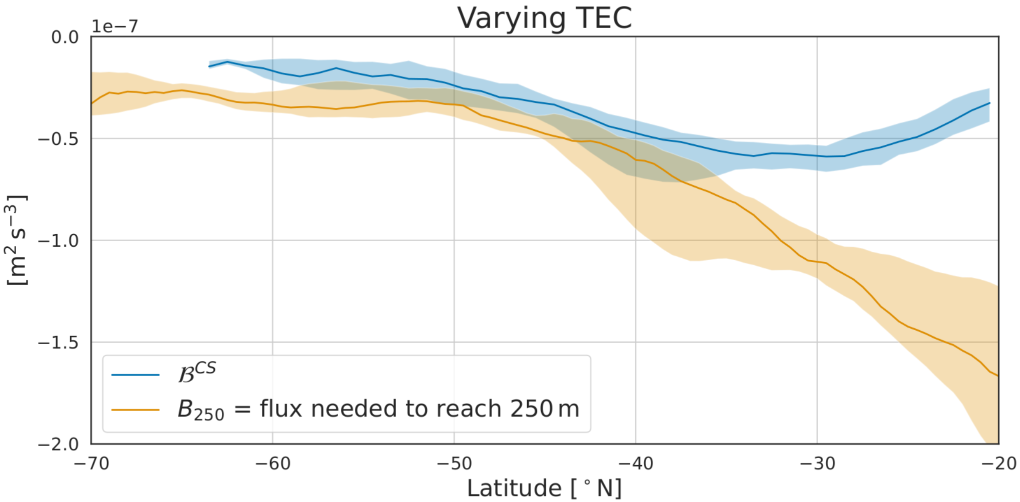
Winter buoyancy loss erodes stratification: zonal view (Paper II) Obj. B



Objective C

Objective A	Objective B	Objective C
Describe alpha – beta oceans using observations	How do buoyancy fluxes shape the upper stratification?	Assess the role of the local value of the TEC.
Paper III	Papers I, II	Papers I, II, and IV

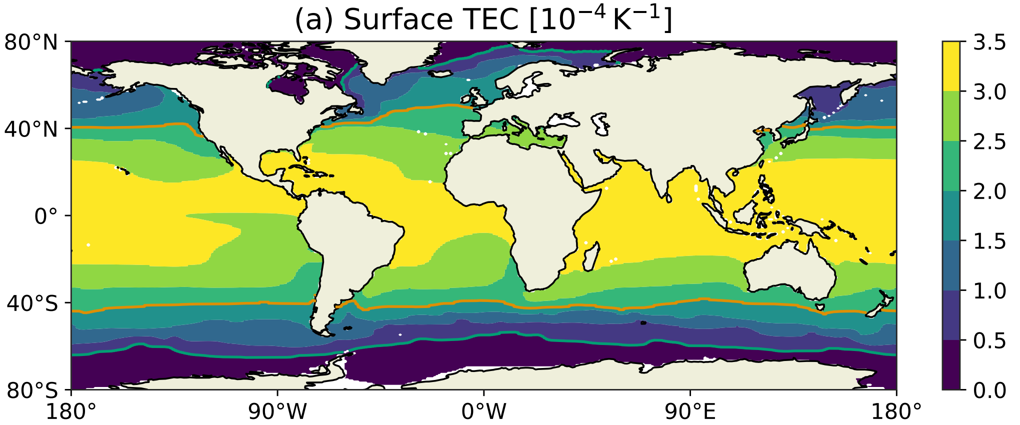
Paper I — Caneill, R., Roquet, F., Madec, G., & Nycander, J. (2022). The Polar Transition from Alpha to Beta Regions Set by a Surface Buoyancy Flux Inversion. *Journal of Physical Oceanography*

Paper II — Caneill, R., Roquet, F., & Nycander, J. (2023). Southern Ocean deep mixing band emerges from competition between winter buoyancy loss and stratification. *accepted for publication in Ocean Science*

Paper IV — Roquet, F., Ferreira, D., Caneill, R., Schlesinger, D., & Madec, G. (2022). Unique thermal expansion properties of water key to the formation of sea ice on Earth. *Science Advances*

The TEC varies with temperature (Paper IV)

- Follows a (quasi) linear relation with temperature
- Decreases the impact of temperature and heat in polar regions



Why does the TEC play a role?

The TEC scales the effect of

temperature on stratification

$$B_{250} = \underbrace{\frac{g}{\Delta t}\int_{-250}^{0} \boldsymbol{\alpha}(z)\frac{\partial \Theta}{\partial z} z \mathrm{d}z}_{B_{250}^{\Theta}} - \underbrace{\frac{g}{\Delta t}\int_{-250}^{0} \beta(z)\frac{\partial S}{\partial z} z \mathrm{d}z}_{B_{250}^{S}} \quad (2)$$

heat fluxes on buoyancy fluxes

$$\mathcal{B}^{surf} = \underbrace{\boldsymbol{\alpha}\frac{g}{\rho_0 C_p} Q_{tot}}_{\mathcal{B}_{\Theta}^{surf}} - \underbrace{\frac{g\beta S}{\rho_0}(E - P - R)}_{\mathcal{B}_{S}^{surf}} \quad (3)$$

$\boldsymbol{\alpha}$ is the TEC

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temperature on stratification

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heat fluxes on buoyancy fluxes

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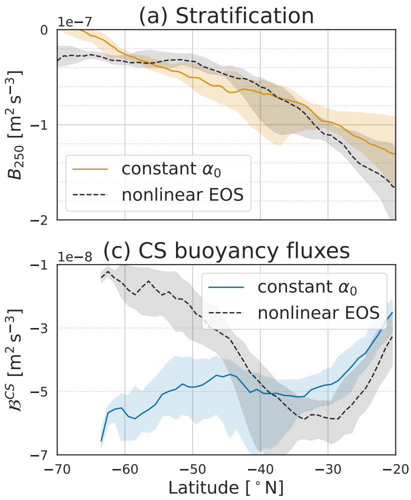
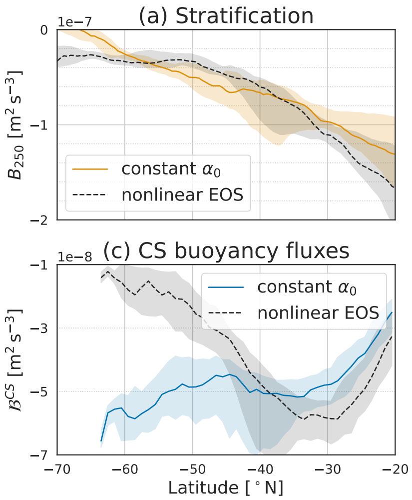
$\boldsymbol{\alpha}$ is the TEC

The impact of the variable TEC (Paper II)



The decrease in the TEC:

- allows for stable beta ocean
- damps buoyancy loss in polar region



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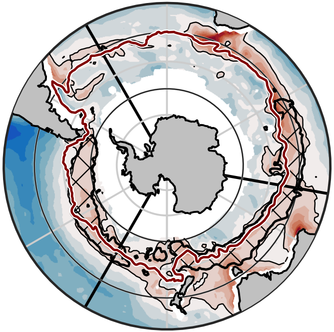
The impact of the variable TEC (Paper II)

$B_{250} - \mathcal{B}^{CS}$

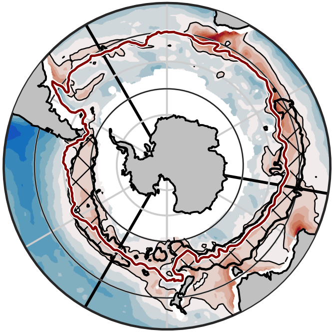


⟵ variable TEC

- The variable TEC controls the width of the DMB
- The decrease in the TEC limits the southward extent of the DMB
- Beta oceans exist because the TEC becomes small

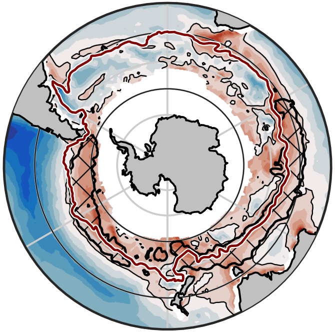
The impact of the variable TEC (Paper II)

$B_{250} - \mathcal{B}^{CS}$



$\longleftarrow$ variable TEC

constant TEC α_0 $\longrightarrow$



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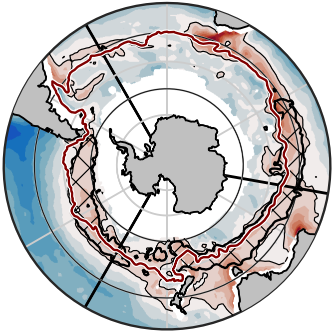

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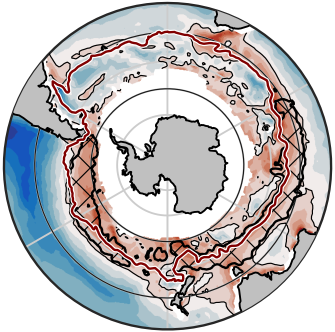

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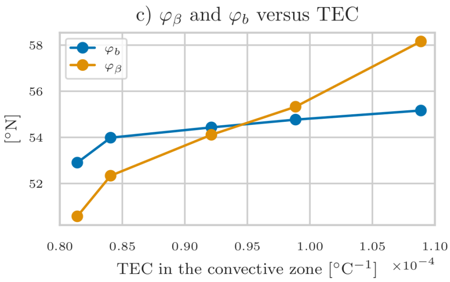
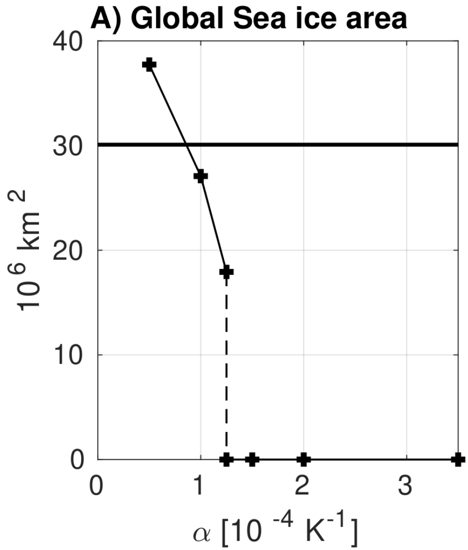
The polar value of the TEC as global controller (Papers II and IV)





Conclusions

Describe alpha – beta oceans using observations. Obj. A

- Global zonation: alpha $\longrightarrow$ transition zone $\longrightarrow$ beta
- ML deeper in alpha- than beta-oceans

How do buoyancy fluxes shape the upper stratification? Obj. B

- The transition zone is located at the sign inversion of annual buoyancy fluxes
- Buoyancy loss erodes stratification and produces the DMB

Assess the role of the local value of the TEC. Obj. C

- The decrease in the TEC in polar regions decreases buoyancy loss
- The small polar value of the TEC permits beta ocean formation
- My thesis confirms that the origin of alpha – beta oceans lies in thermodynamic of seawater

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- The decrease in the TEC in polar regions decreases buoyancy loss
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Perspectives

- The sea surface temperature exerts a strong control on the stratification by its link with TEC.
- Buoyancy fluxes are not simply the sum of heat and freshwater fluxes.
- Warming $\implies$ larger values of the TEC. But also increases freshwater fluxes in the polar regions. Who will win?

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